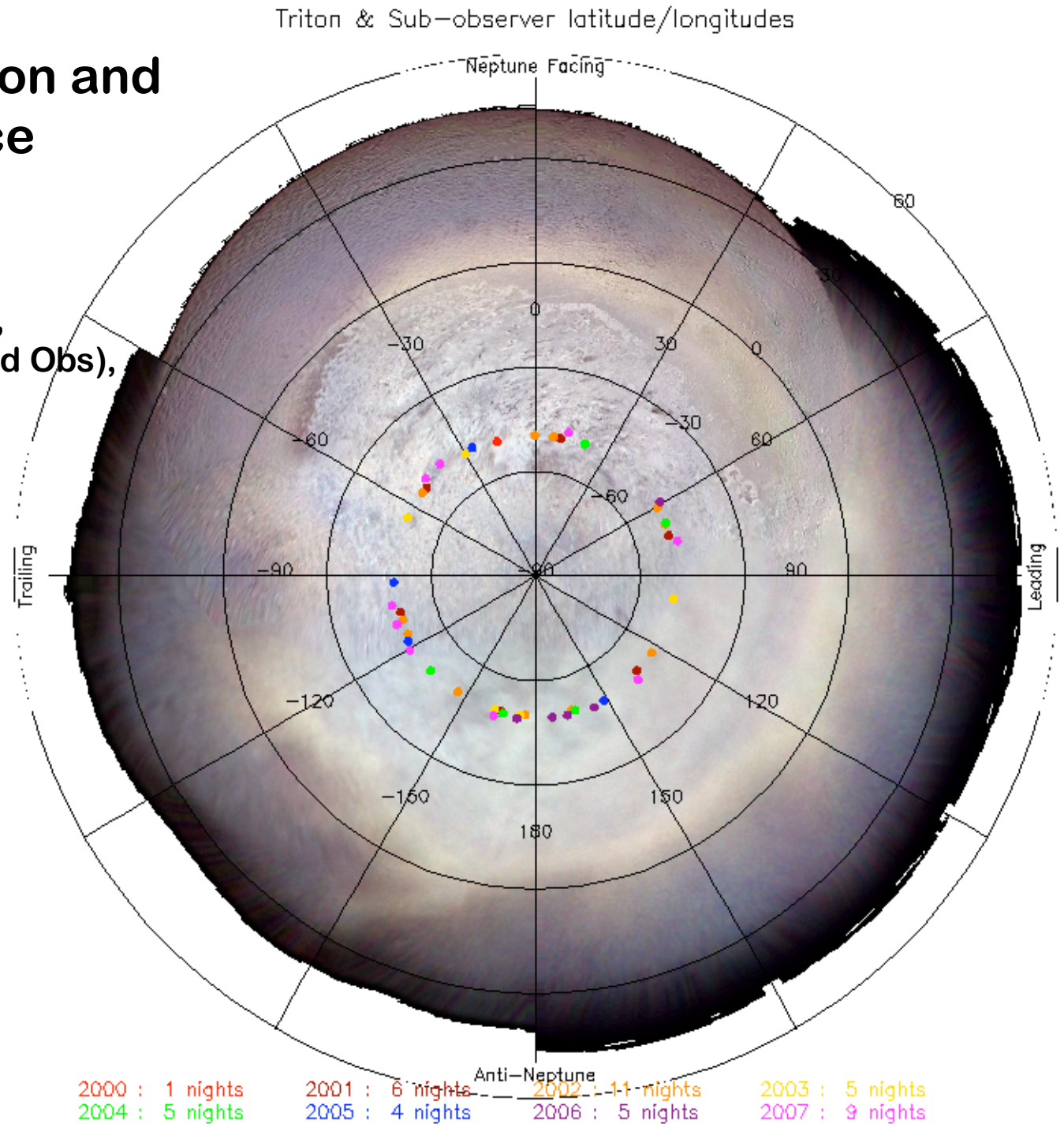
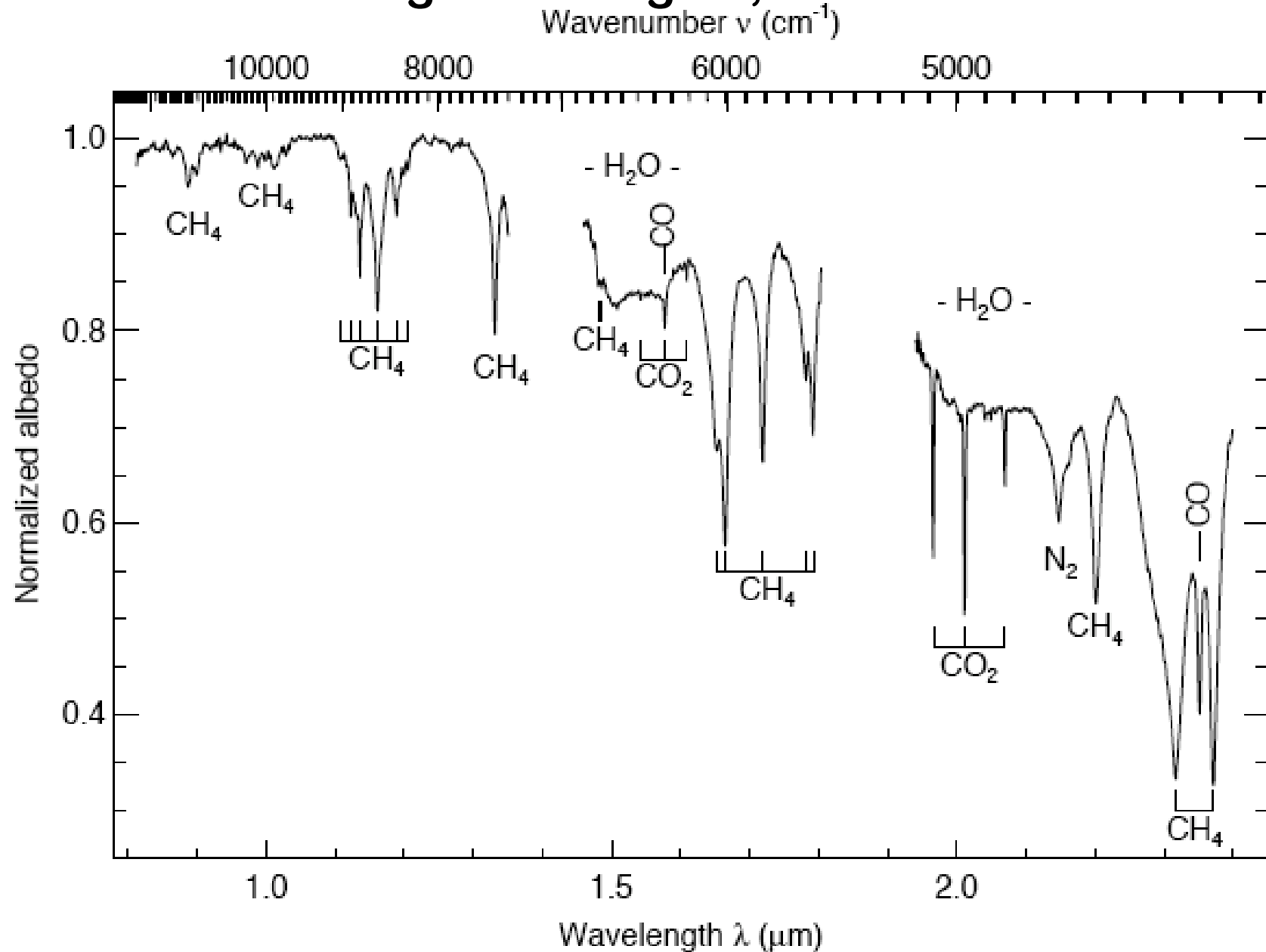


Spatial distribution and evolution of N₂ ice on Triton

Leslie Young (SwRI),
Will Grundy (Lowell Obs),
John Stansberry (Steward Obs),
Marc Buie (SwRI)



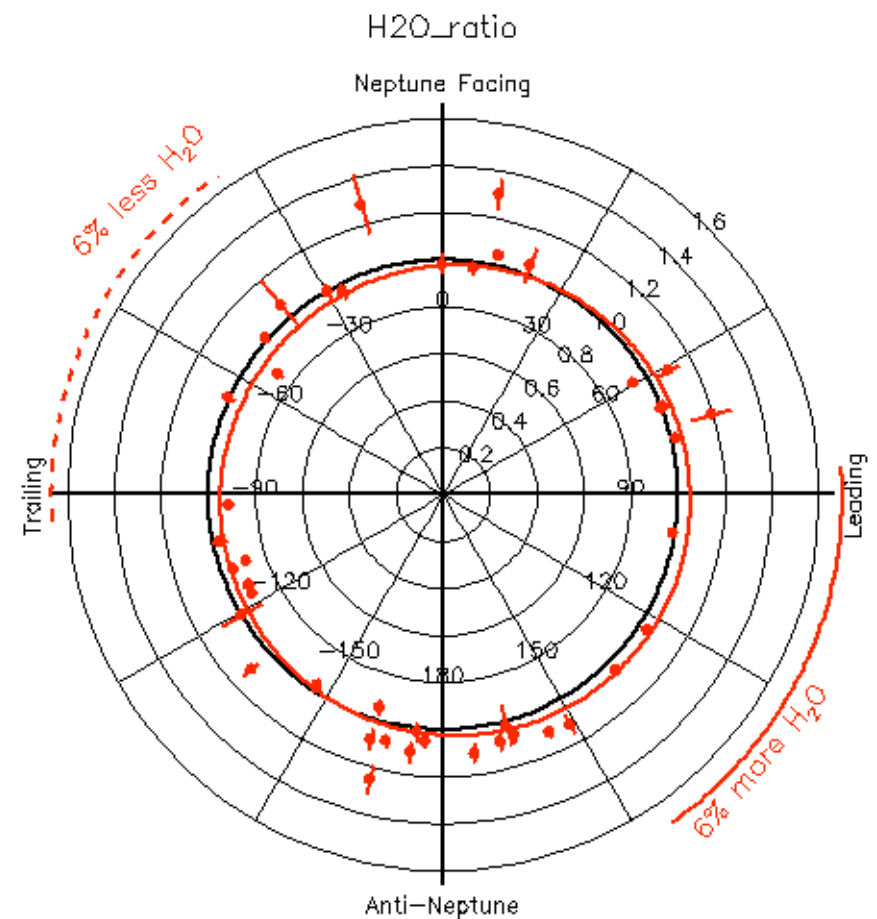
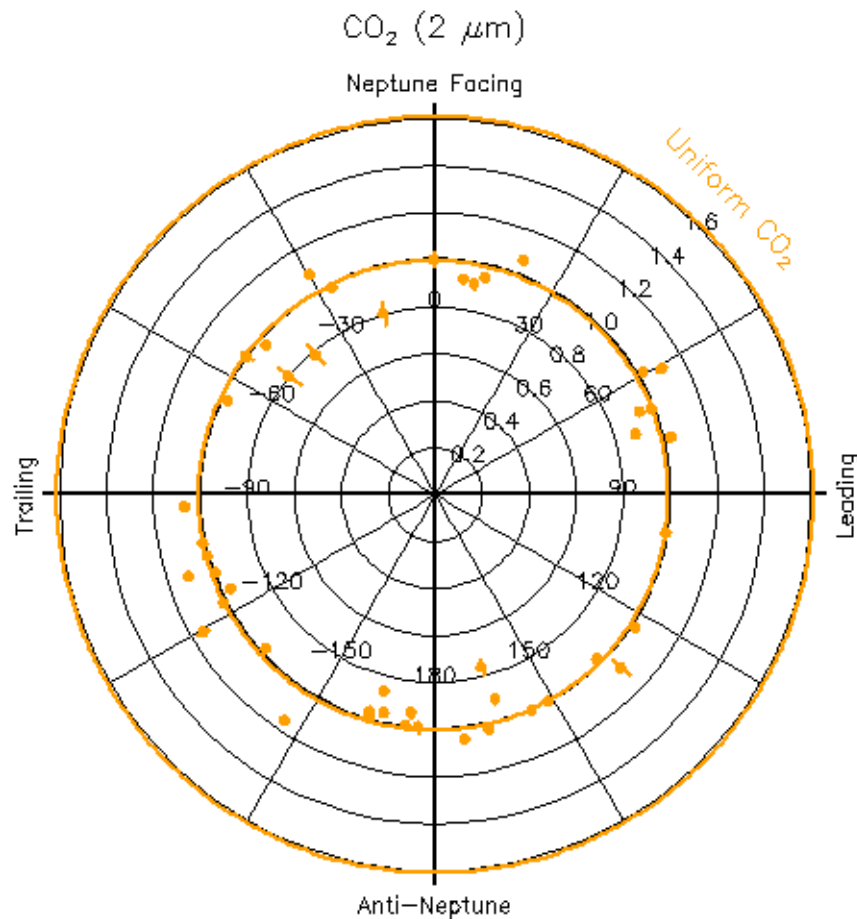
Wavenumber ν (cm^{-1})

The bedrock: H₂O and CO₂

H₂O and CO₂ are non-volatile at Triton's temperatures.

H₂O shows a small variation with longitude ($11 \pm 2\%$ peak-to-peak), while

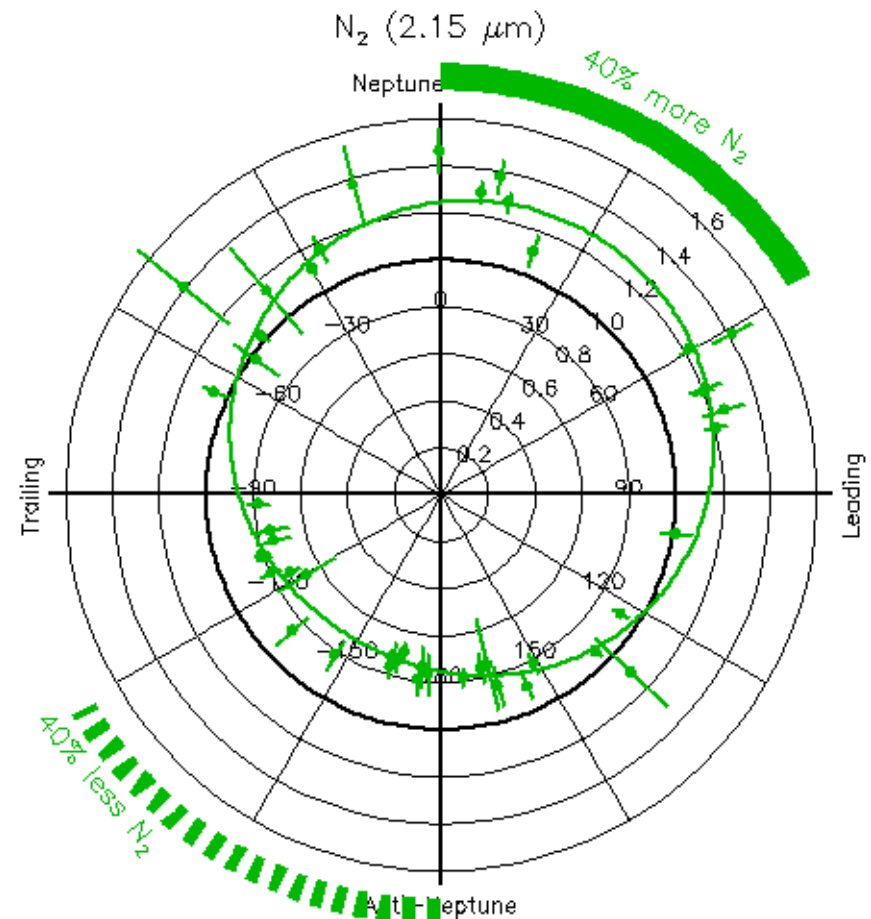
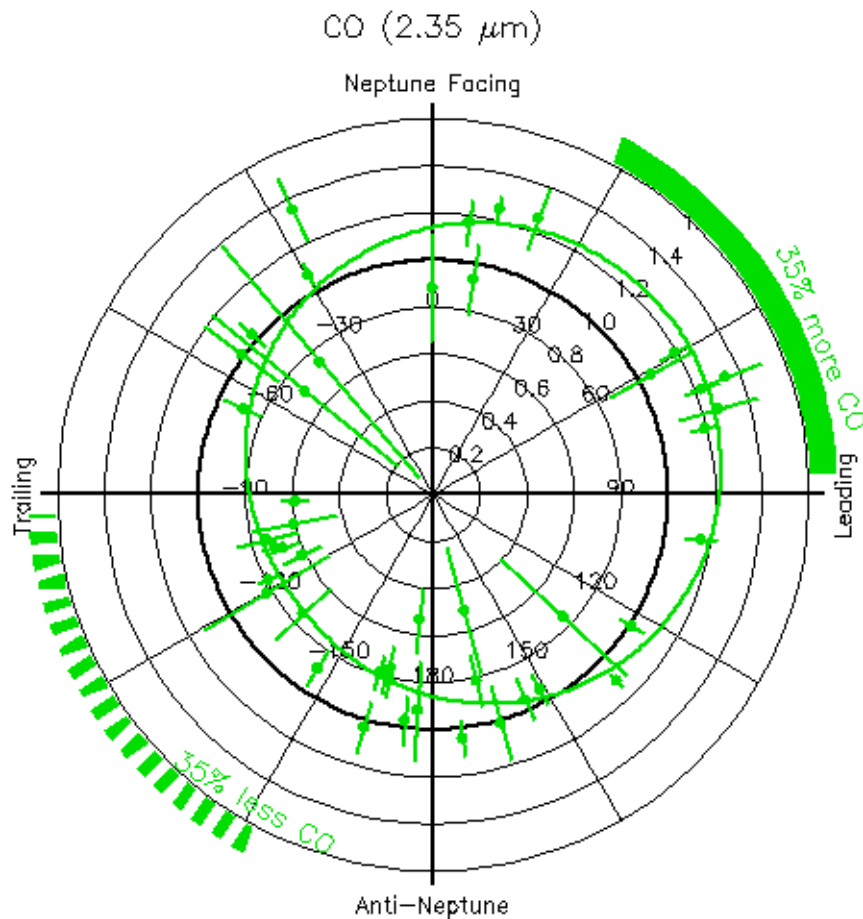
CO₂ shows a no variation ($2 \pm 2\%$ peak-to-peak)



The highly volatile ices: N_2 and CO

N_2 is the most volatile species on Triton's surface, and so it is the dominant species in Triton's atmosphere, and details of emissivity, the N_2 ice distribution, and albedo controls Triton's surface pressure.

CO and N_2 can mix at any ratio, and their vapor pressures are similar, so it is not surprising to see them vary together.

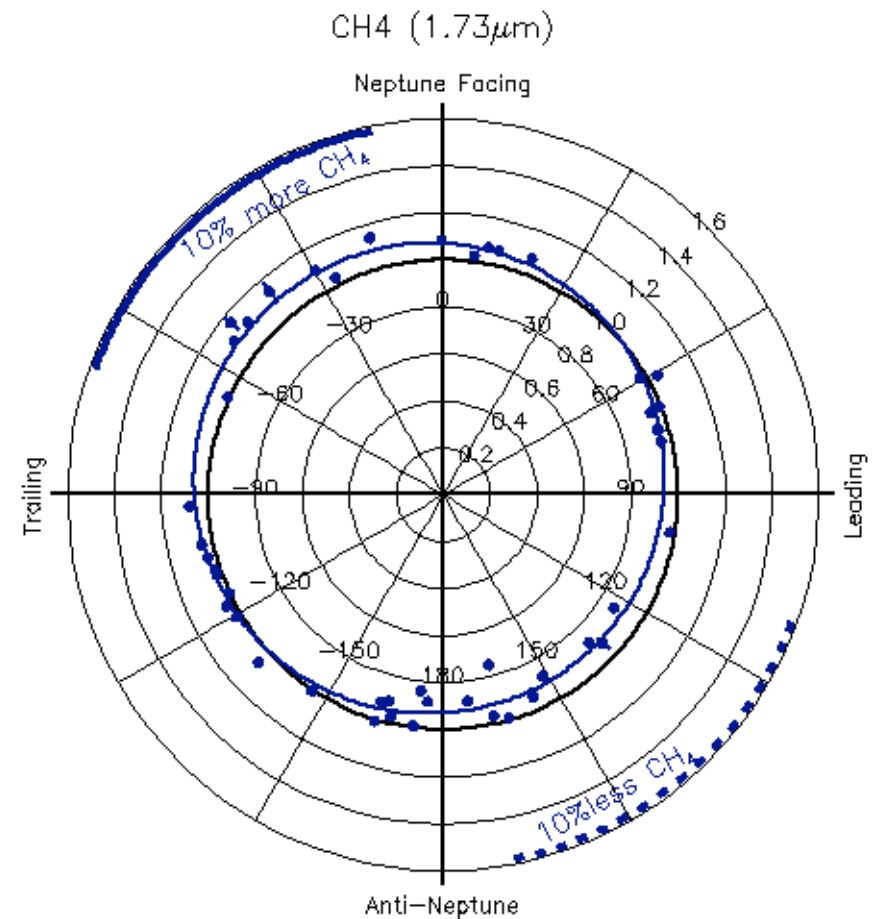
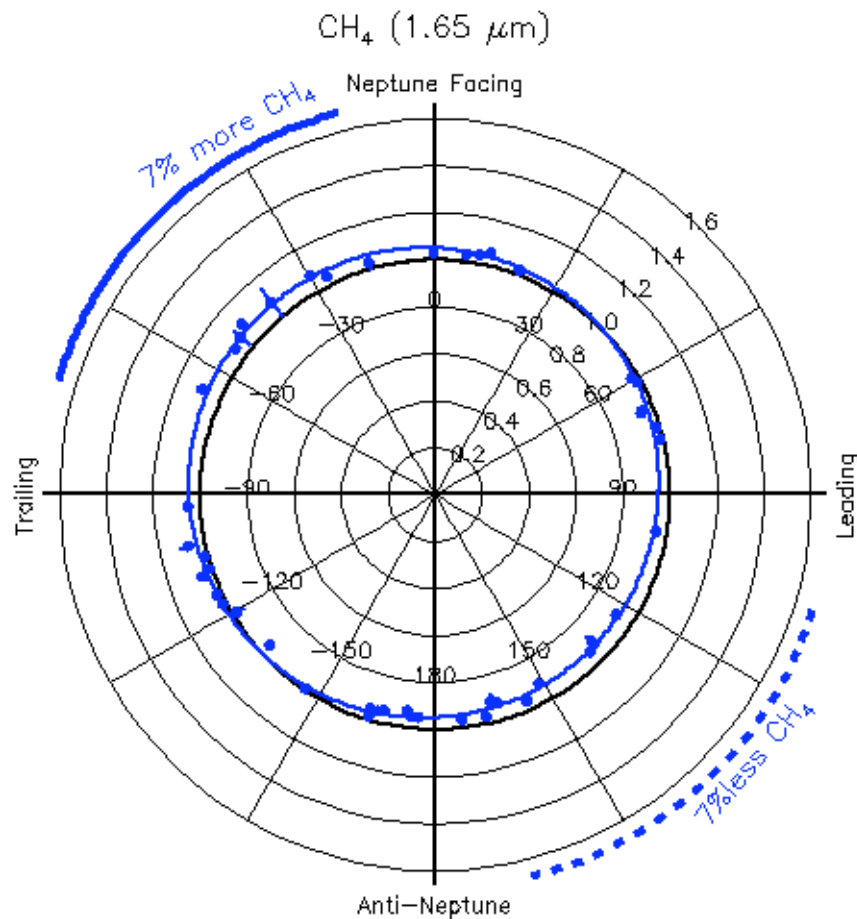


Widespread CH₄ (from strong bands)

Strong bands of CH₄ (e.g., 1.65, 1.73 μm) are sensitive to shallow films or to dilute amounts, as well as to higher concentrations.

This widespread CH₄ shows some variation, depending on the band ($\sim 17 \pm 2\%$ peak-to-peak)

The CH₄ does *not* peak at the same longitudes as N₂, so the N₂:CH₄ mixing ratio must vary with location.

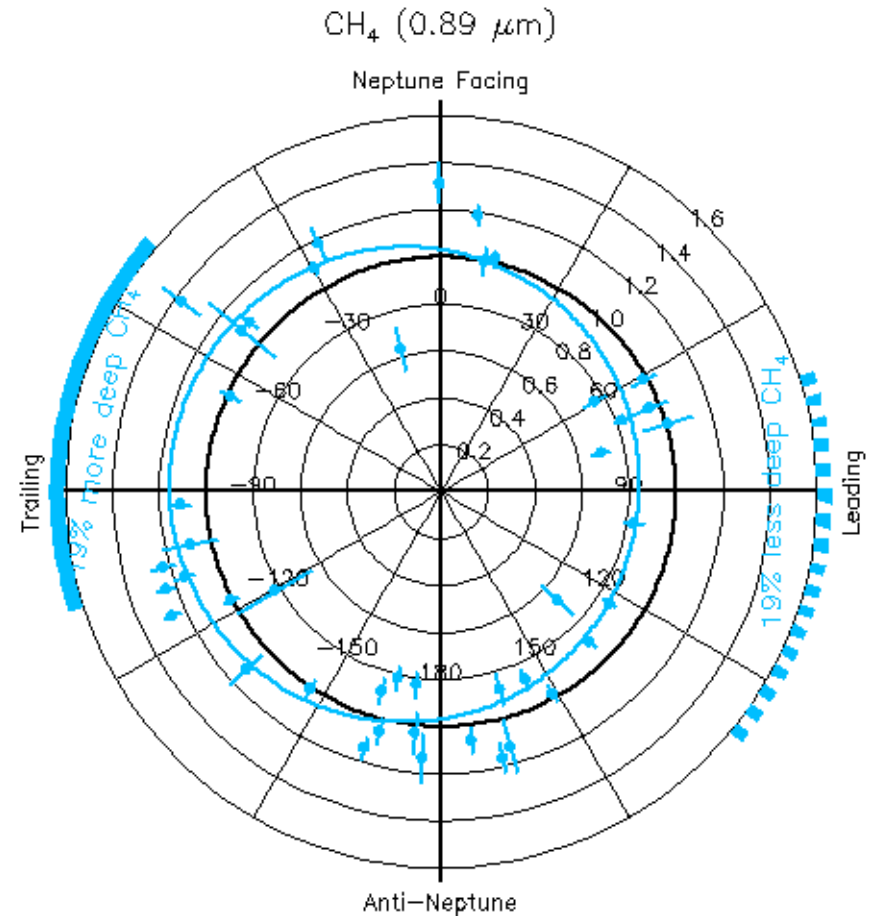
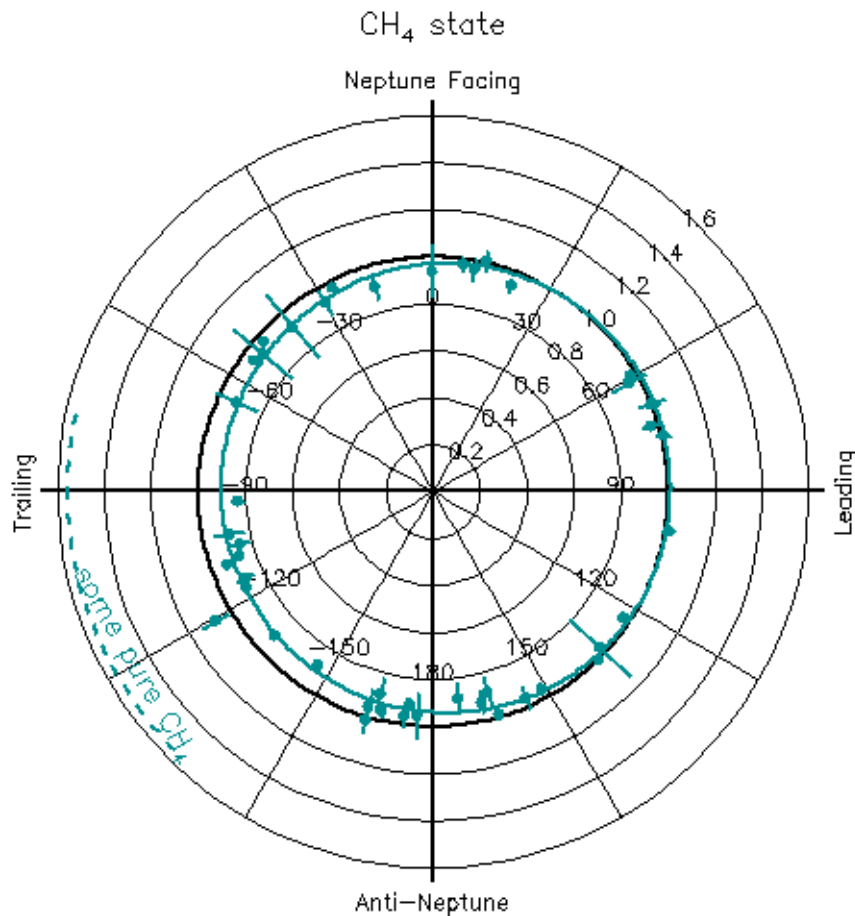


Localized relatively concentrated CH₄ (from weak bands & band shifts)

Weak bands of CH₄ (e.g., 0.89 μm) are sensitive to deeper layers or to concentrated amounts.

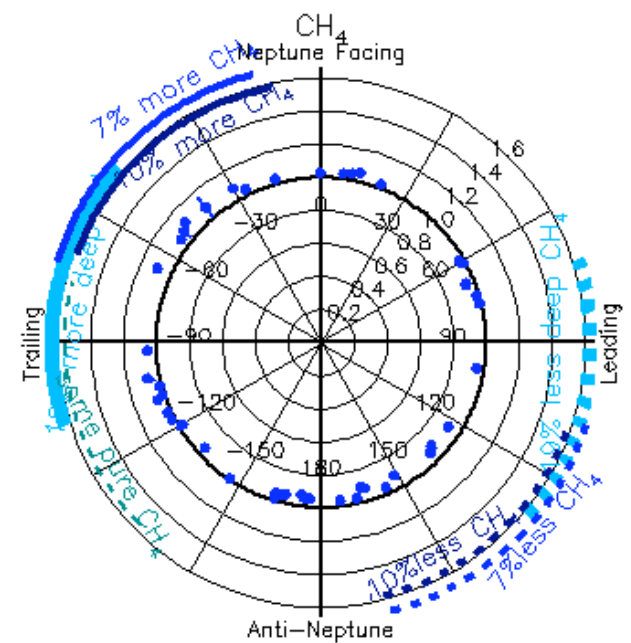
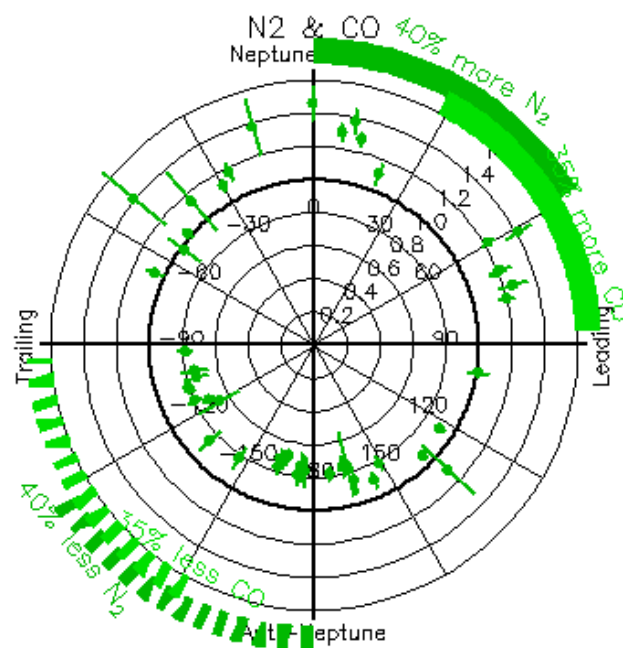
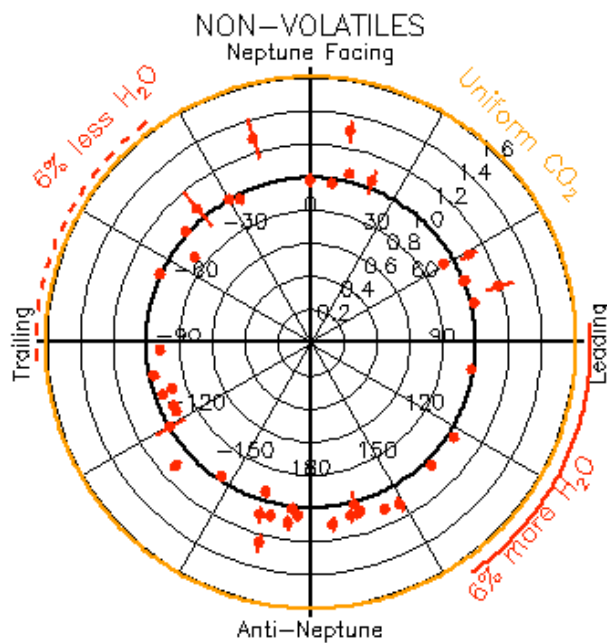
This "deep" CH₄ shows more variation than the "shallow" CH₄
($\sim 38 \pm 6\%$ peak-to-peak, vs. $15 \pm 1\%$ for the 1.65 μm band)

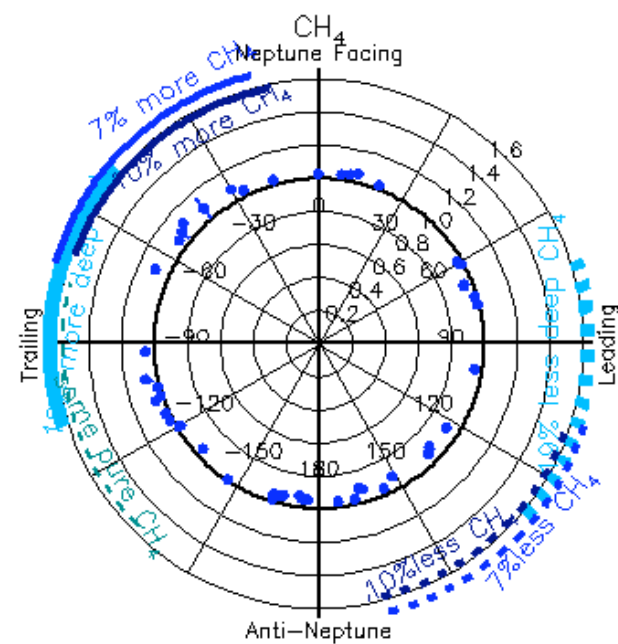
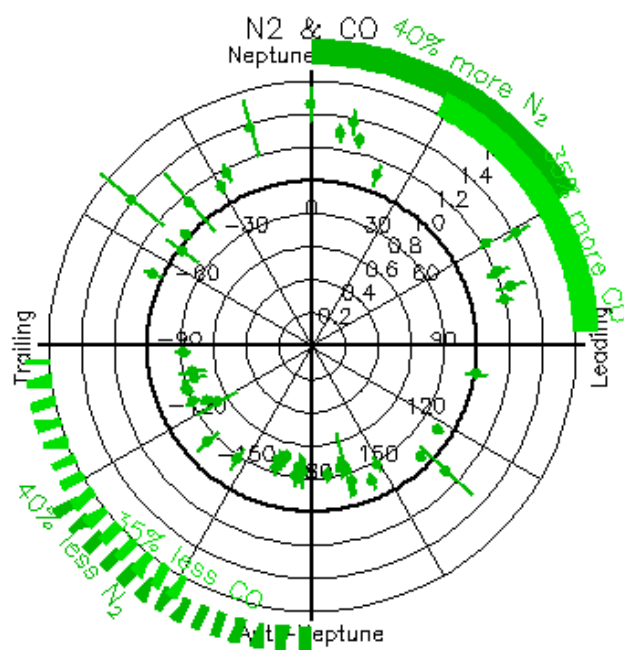
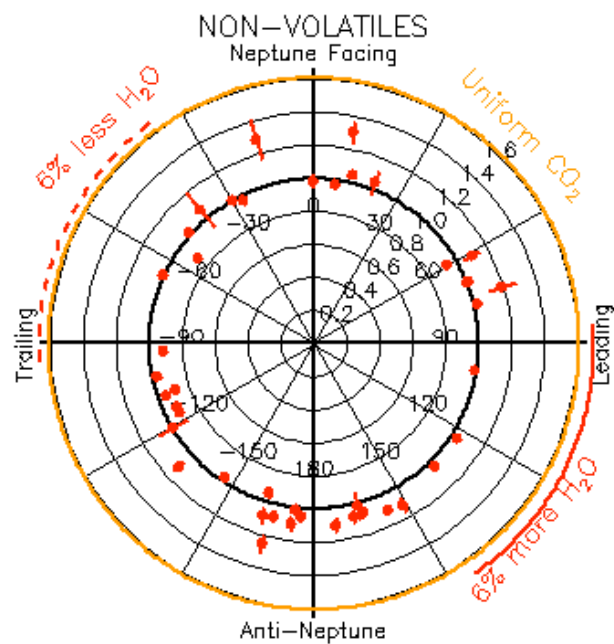
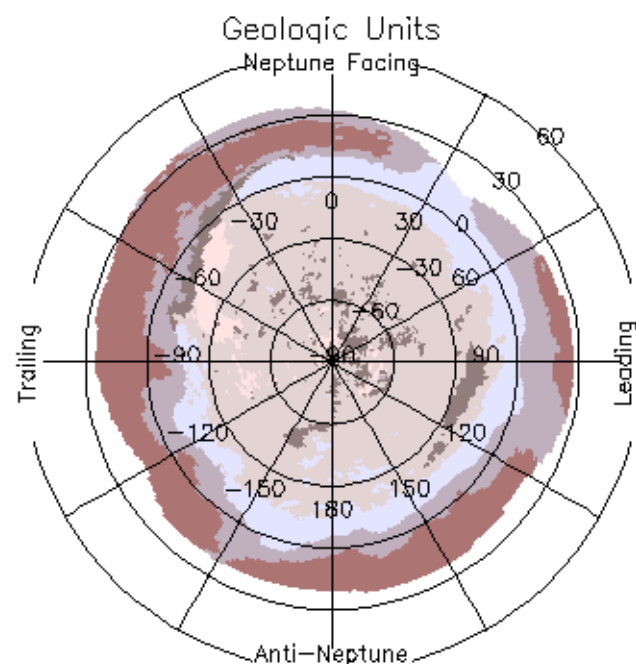
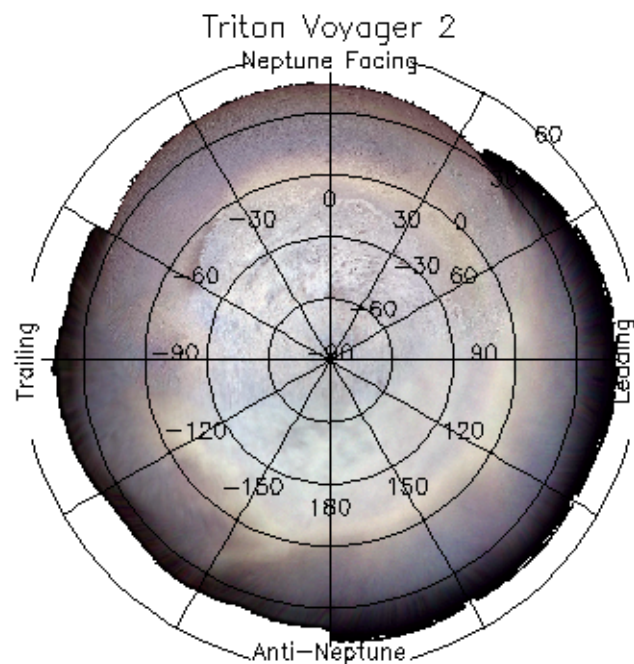
CH₄ lines shift with wavelength when CH₄ is diluted in N₂. Nearly all of the CH₄ on Triton is diluted in N₂, but some pure CH₄ (or CH₄ with trace N₂) apparently exists near -120°

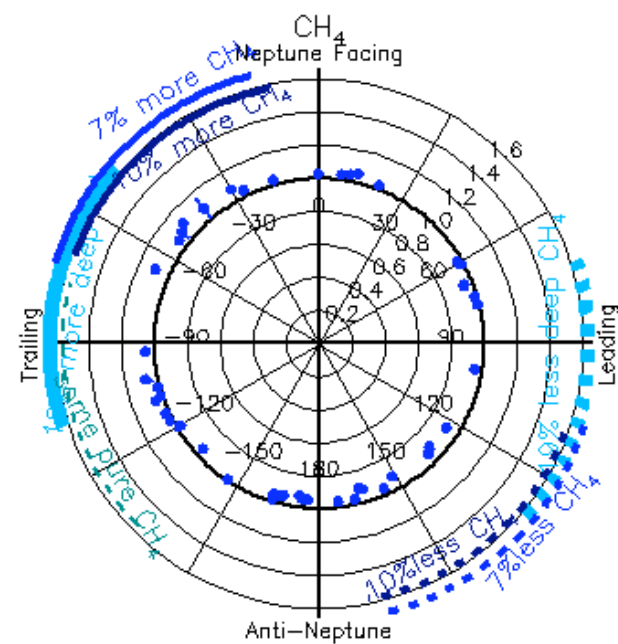
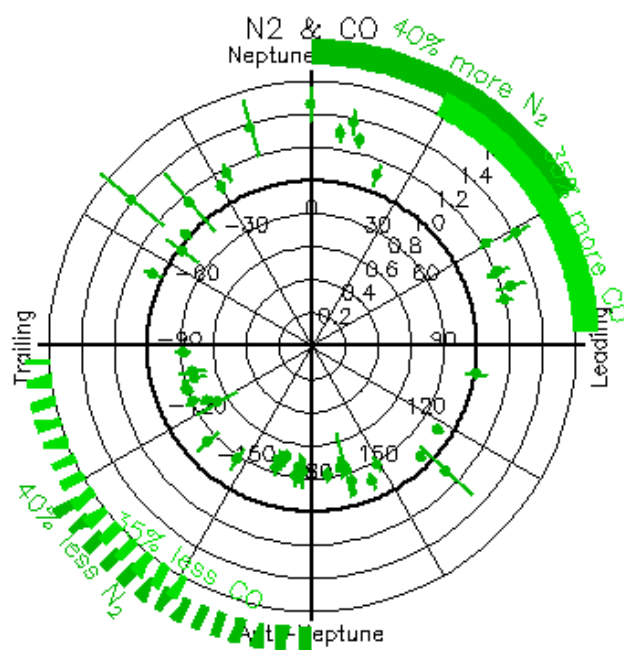
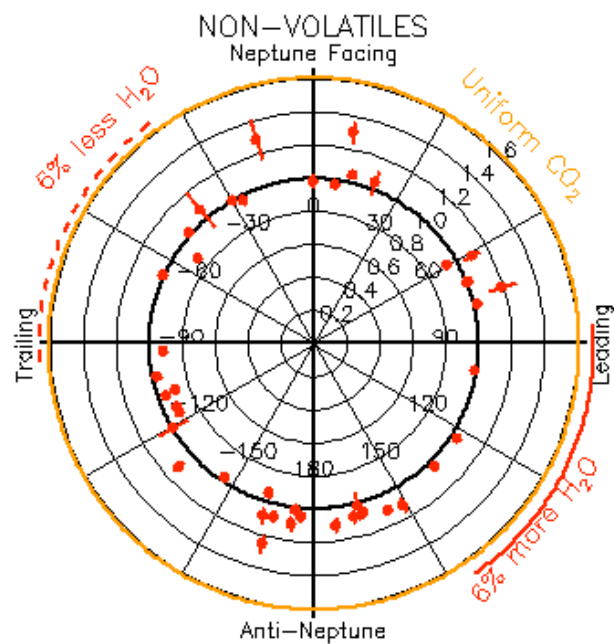
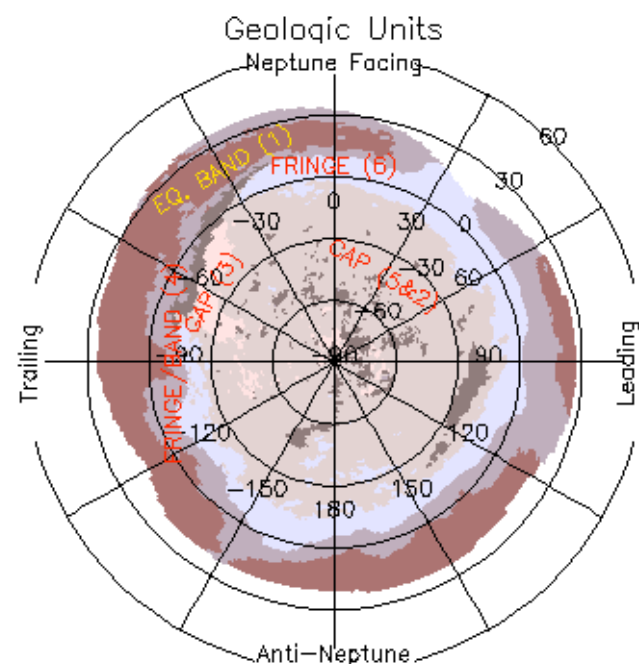
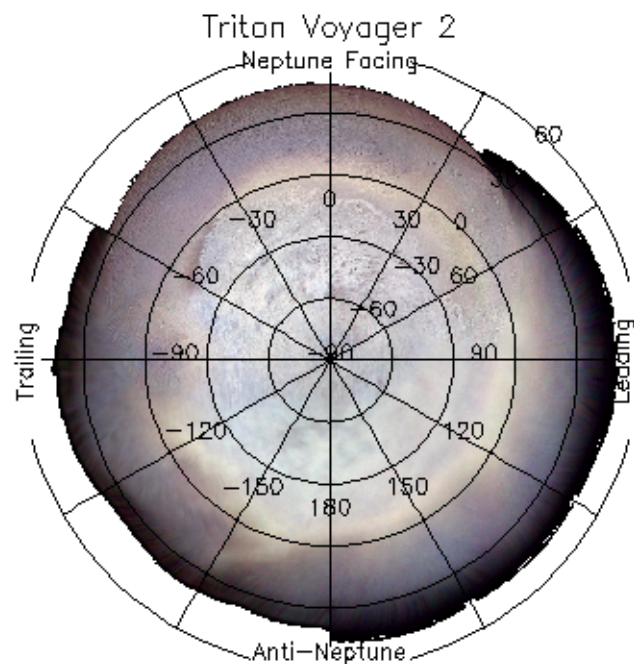


Summary of longitudinal correlations

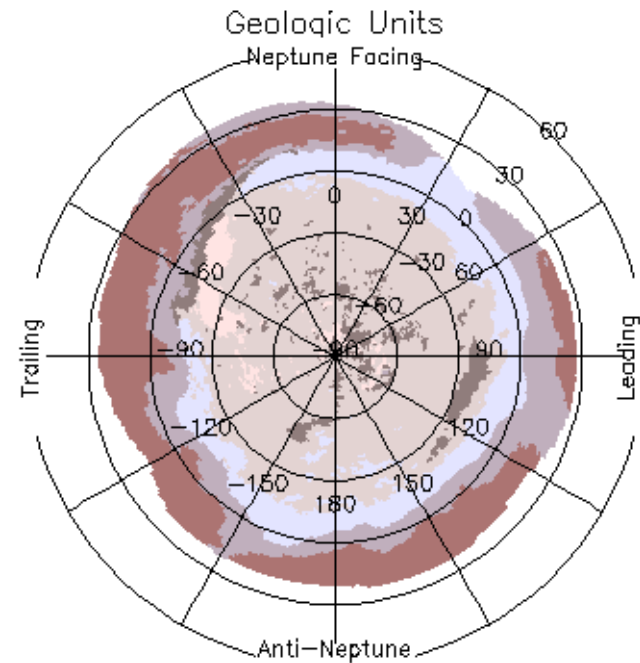
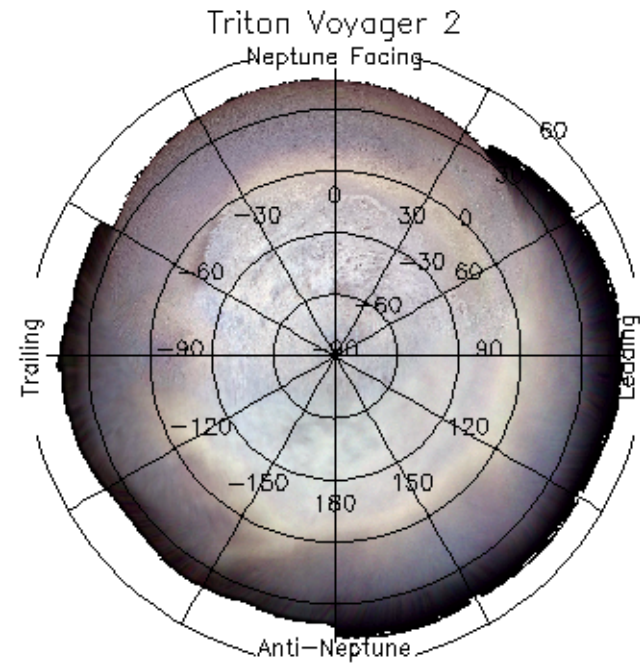
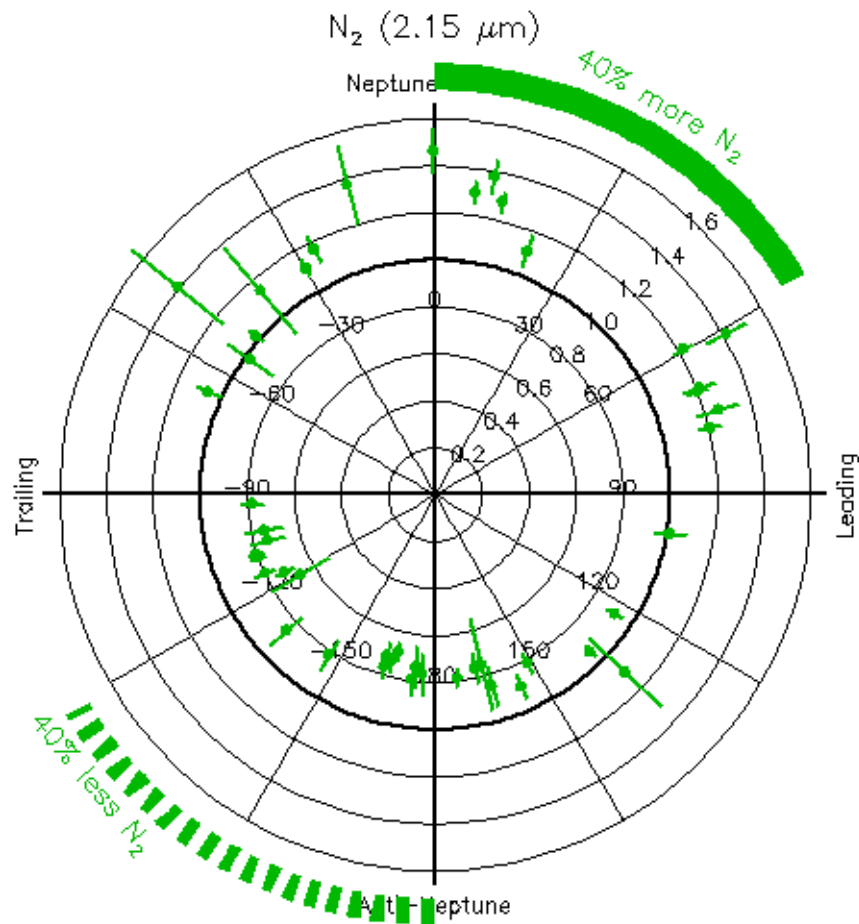
- N_2 and CO are strongly correlated
- H_2O is weakly anticorrelated with wide-spread CH_4
- CH_4 is *not* correlated with N_2
- Purer CH_4 is weakly anti-correlated with N_2







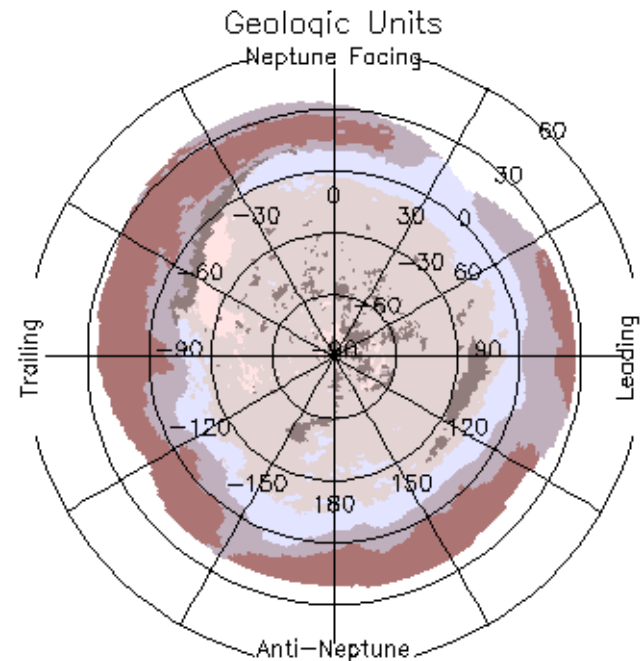
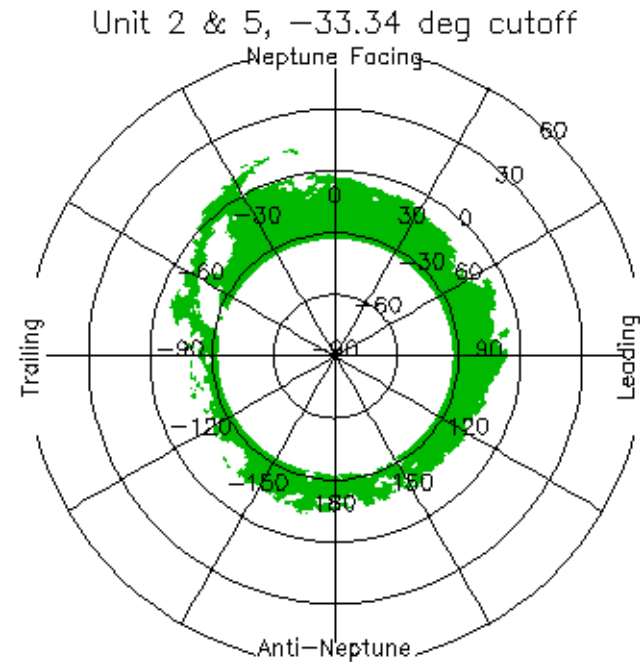
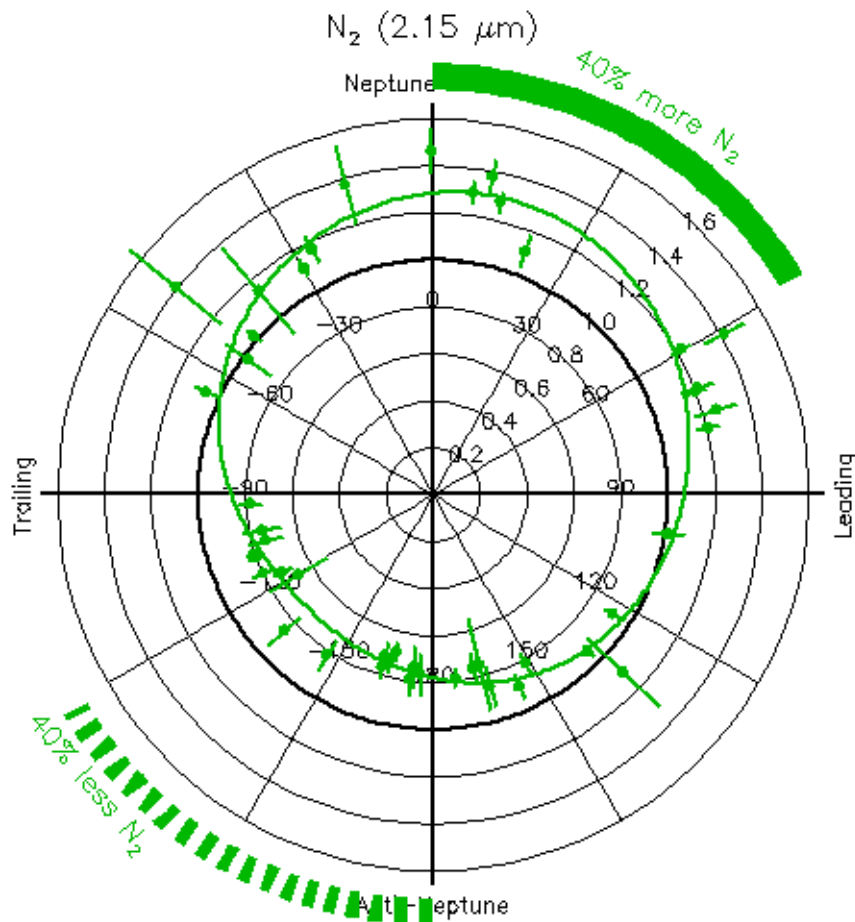
The N_2 variation doesn't directly relate to the geologic units identified by McEwen et al. 1990.



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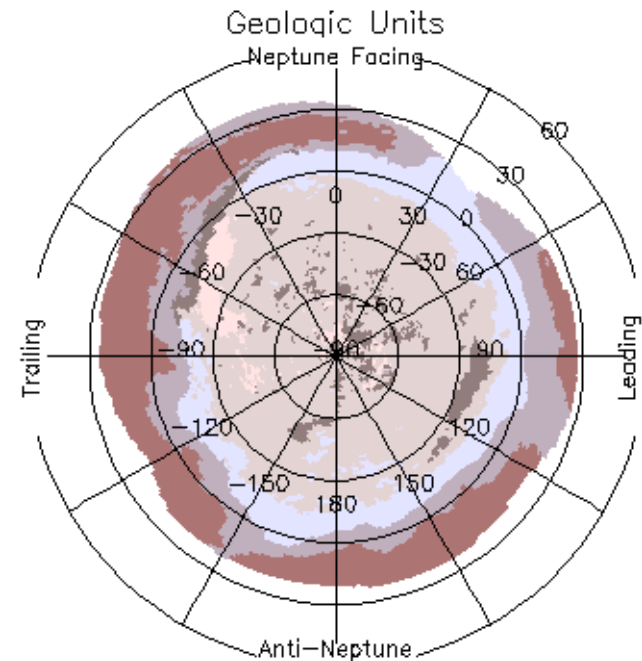
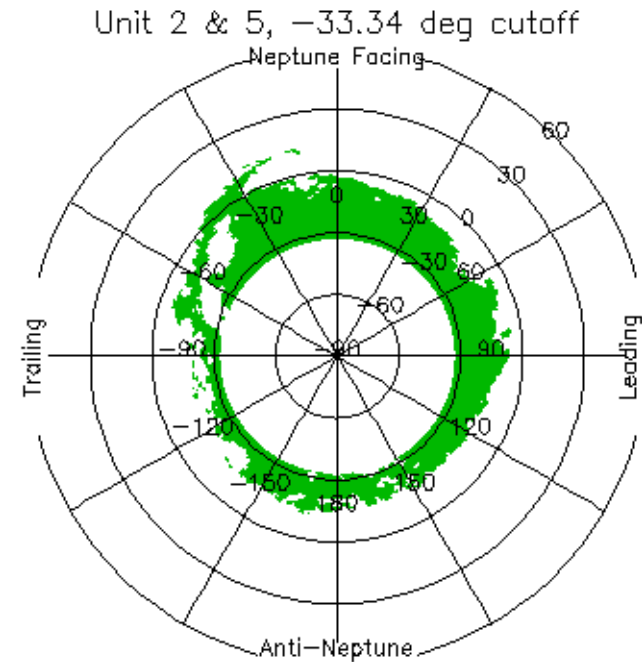
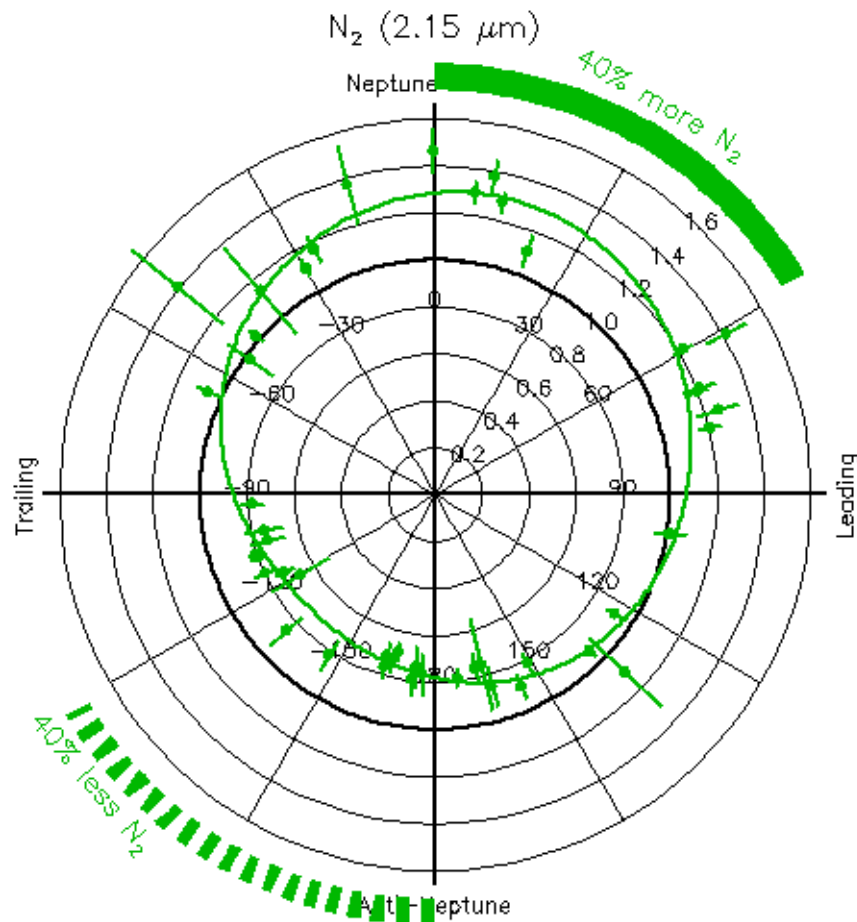
Excellent matches to the N_2 variation can be made from the cap units with a denuded pole.

Other models work well too: including the fringe or not, or assuming detectable N_2 only where diurnally averaged insolation is below a threshold.



But consider the atmosphere...

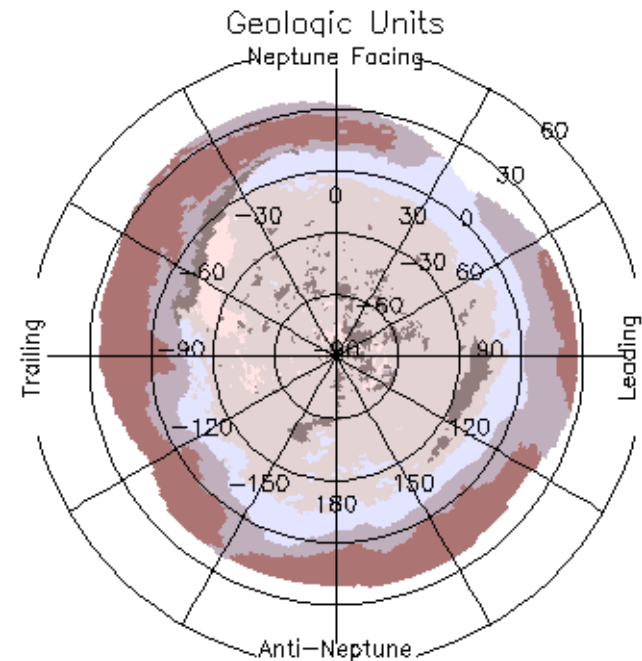
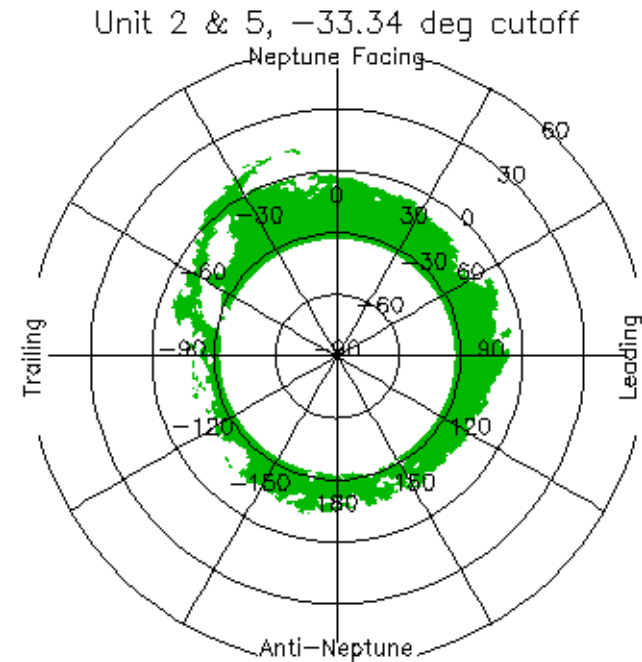
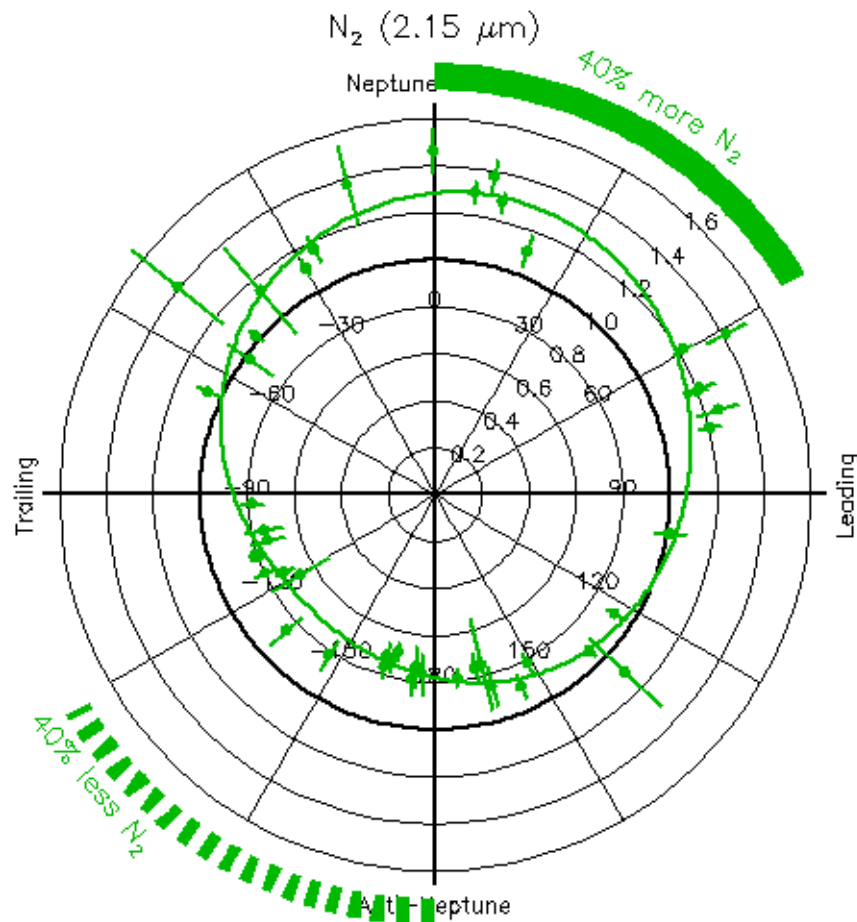
- If the N_2 is on the winter hemisphere, how can the emissivity be low enough (~ 0.2) to match Triton's surface pressure (corresponding to a N_2 ice temperature of $\sim 38K$)?
- If the N_2 is *not* at the winter pole, why not? Is there a heating affect from an internal heat source?



How much fine-grained N_2 (which is not detectable at $2.15 \mu m$) can we hide on the southern hemisphere?

- The wide-spread CH_4 is diluted in the N_2 argues for "invisible southern N_2 ." (For example: mm-sized at the pole and cm-sized at the edge.)
- The emissivity needed for 38 K is reasonable.

The denuded pole shows no obvious signature of the difference in N_2 texture in the Voyager images.



Food for thought (and requests for lab data)

- How does the detectability of N_2 depend on whether it is a sublimation or condensation site?
- Does the CH_4 spectrum in $N_2:CO:CH_4$ differ from that in $N_2:CH_4$?
- How does the emissivity of N_2 ice depend on CO and CH_4 concentrations, the rate of sublimation or condensation, and age?
- What are the sticking coefficients of N_2 , CO, and CH_4 with respect to the various units?
- How do CO and CH_4 diffuse through N_2 ?